

Evaluating the performance of Italian intensity prediction equations and analysis of directivity using macroseismic data

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SECURE Mid-term - 19-20 Giugno 2025, Milano



Istituto Nazionale di Geofisica e Vulcanologia



SECURE
REGIONAL-SCALE EARTHQUAKE GROUND MOTION PREDICTIONS
THROUGH PHYSICS-BASED AND EMPIRICAL APPROACHES:
A CASE STUDY CENTRAL ITALY

INTRODUCTION

The **macroseismic intensity data** expresses the level of ground shaking generated at a given site by an earthquake

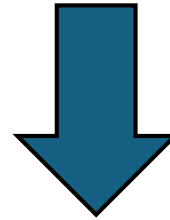


In Italy a large amount of macroseismic data that has greatly improved the **Italian Macroseismic Database – DBMI** and the **Italian Parametric Earthquake Catalogue - CPTI**

This allowed the development of several **Intensity Prediction Equations (IPEs)** and their continuous updating

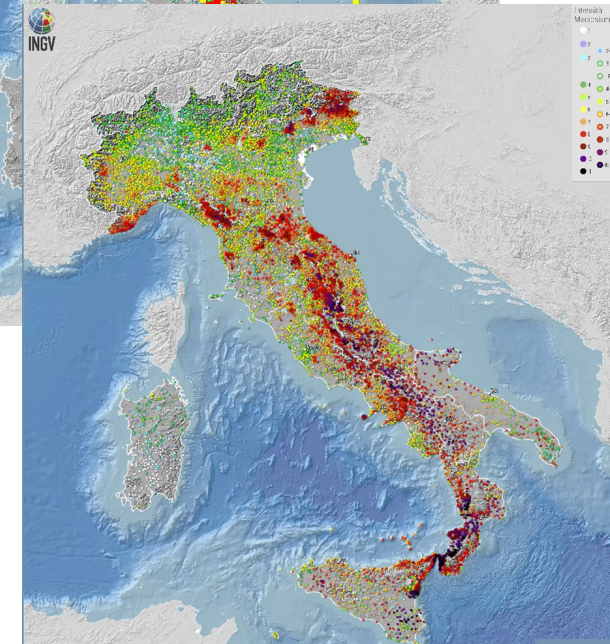
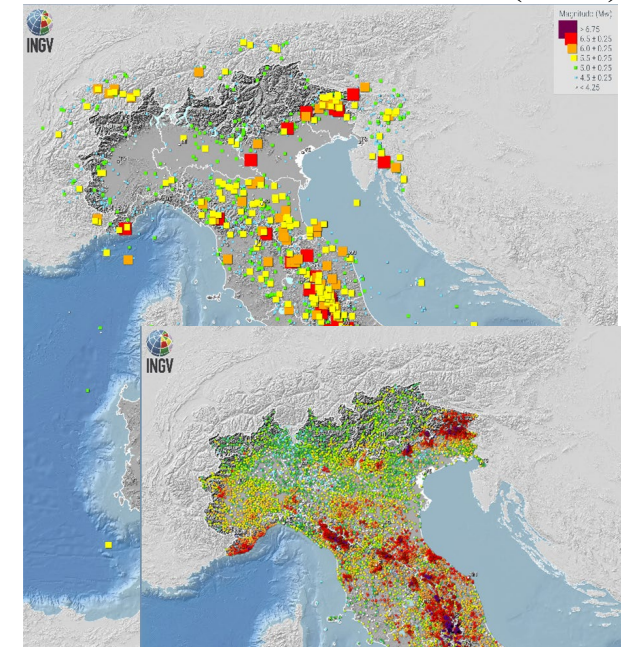
GOAL

We evaluated the performance of the Italian attenuation models by comparing **the documented intensity data** of **DBMI15** with **the predicted values** computed from the epicentral parameters collected in the **CPTI15**



1. We selected the **most recent IPEs** valid for the Italian area
2. We considered **four datasets** derived from the last version of CPTI15 and DBMI15
3. We **predicted the intensity value** at a given site of the selected datasets
4. We analyzed the results to define the **best predictive model** following the GMMs strategy (Lanzano et al., 2020)

Rovida et al. (2022)



Locati et al. (2022)

1) IPEs selection

We selected 5 intensity attenuation

$$D = \sqrt{R^2 + h^2}$$

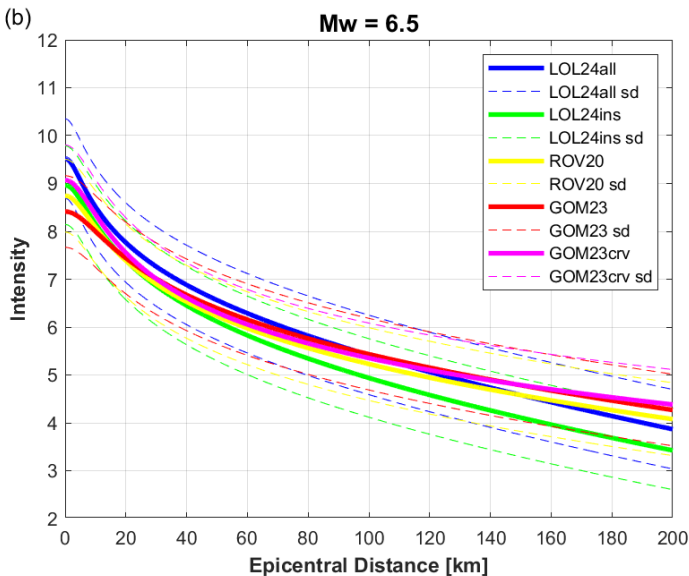
ID	Reference
ROV20	Rovida et al. 2020
LOL24all	Lolli et al. 2024
LOL24ins	Lolli et al. 2024
GOM24	Gomez-Capera et al. 2024
GOM24crv	Gomez-Capera et al. 2024

$$I = I_E - b * (D - h) - c * [\ln(D) - \ln(h)]$$

$$I_E = a + d * M_w$$

$$I = a - b * D - c * \text{Log}(D) + d * M_w$$

$$\text{Log}I = a - b * D - c * \text{Log}(D) + d * \text{Log}M_w$$



ID	N° Events	Time range	Mw range	Focal Depth	N° IDPs	Int range
ROV20	354	1903–2013	2.7–7.1	<30 km	30,138	2–11
LOL24all	744	1200–2017	*	*	33,038	3–11
LOL24ins	338	1904–2017	*	*	20,029	3–11
GOM24	119	1908–2013	3.8–7.1	<35 km	16,260	3–11
GOM24crv	119	1908–2013	3.8–7.1	<35 km	16,260	3–11

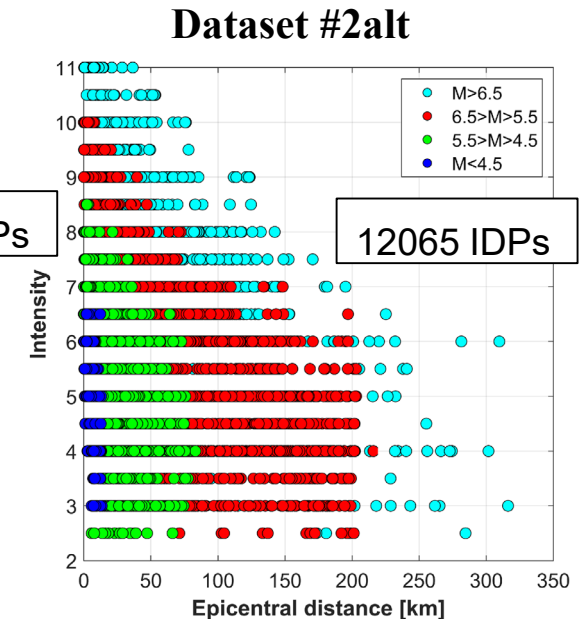
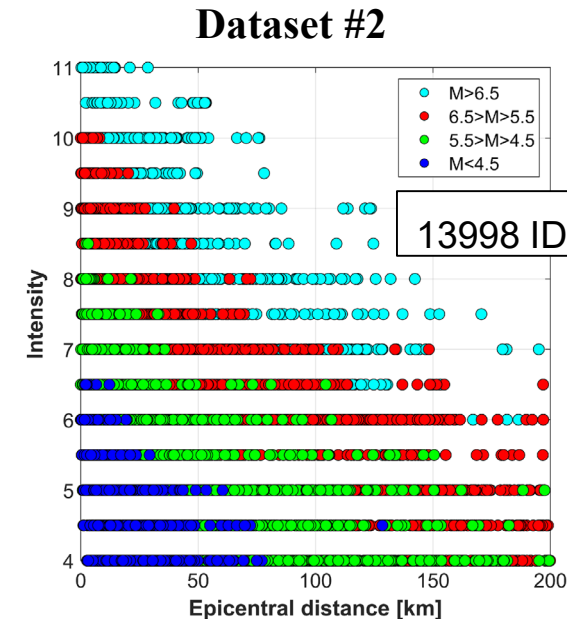
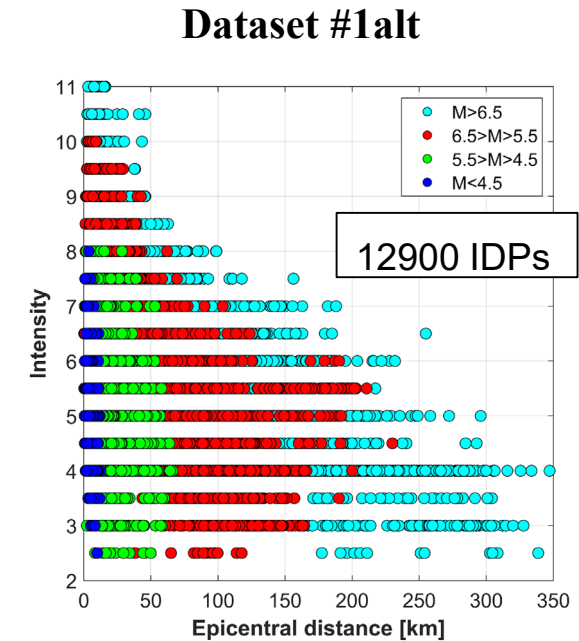
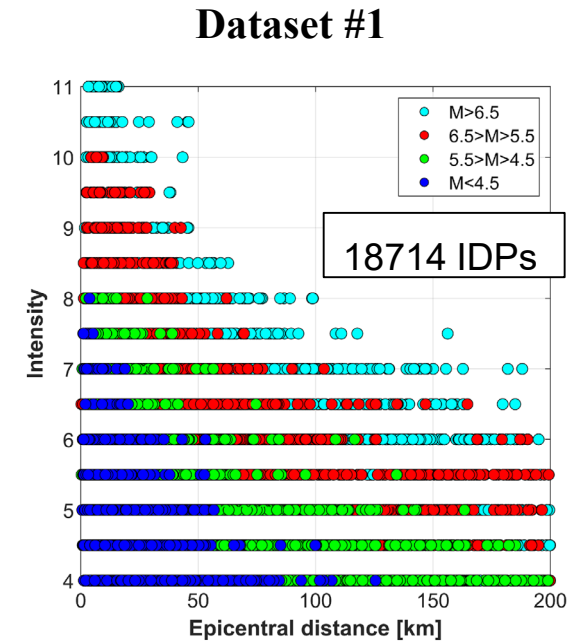
2) Testing Datasets

We selected 4 different datasets:

- Number of associated IDPs ≥ 10
- Moment magnitude $M_w \geq 4.0$
- No volcanic and subduction earthquakes

1. Two datasets with **213 earthquakes** with **instrumental magnitudes and location** from 1907 to 2020

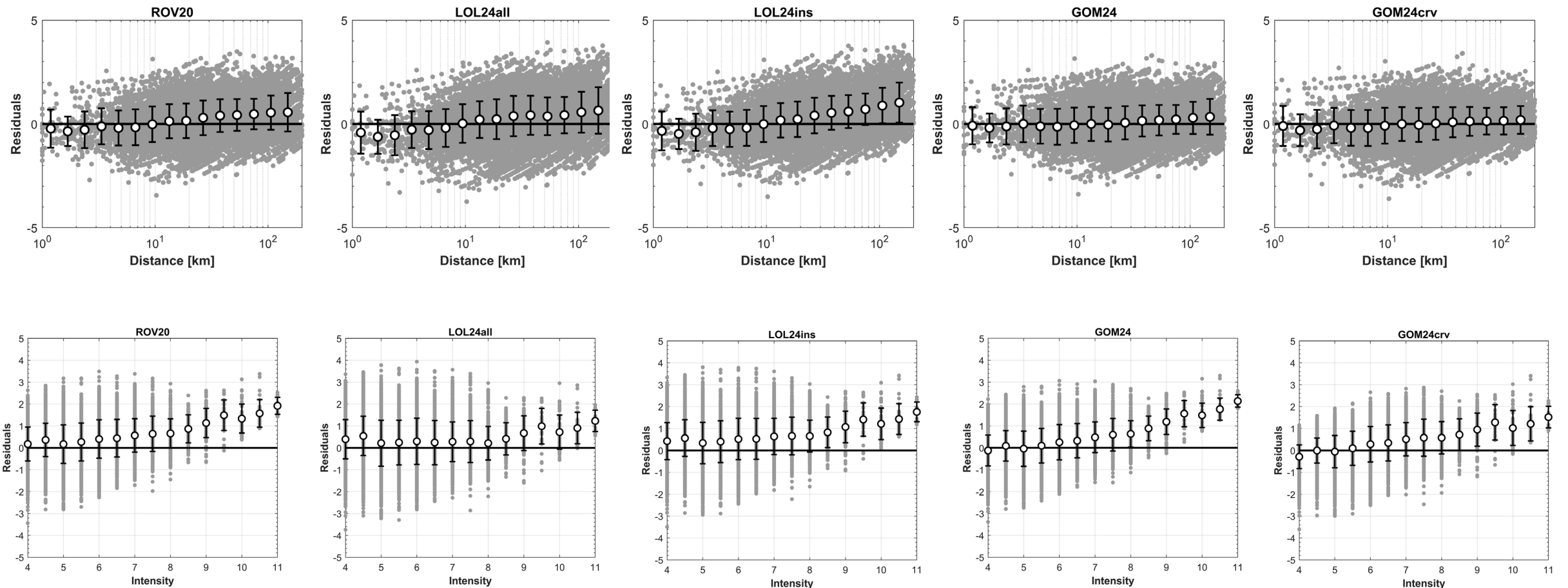
2. Two datasets with **357 earthquakes** with **parameters derived only from macroseismic data** from 1117 to 1978



3) Analysis of the residuals

The total residual *Res* is calculated as the difference between the macroseismic intensity observations *MI_{obs}* (from the selected datasets) and the predictions at the same sites *MI_{pred}*

$$R_{es} = MI_{obs} - MI_{pred}$$



4) Best-performing model

We adopted two metrics to evaluate the performance of the predictive models:

- (i) the **log-likelihood (LLH)** proposed by Scherbaum et al. (2009) in the framework of PSH assessment for Europe to rank both IPEs and GMMs
- (ii) the **mean absolute error (MAE)**, introduced by Mak et al. (2015) for validating Italian IPEs

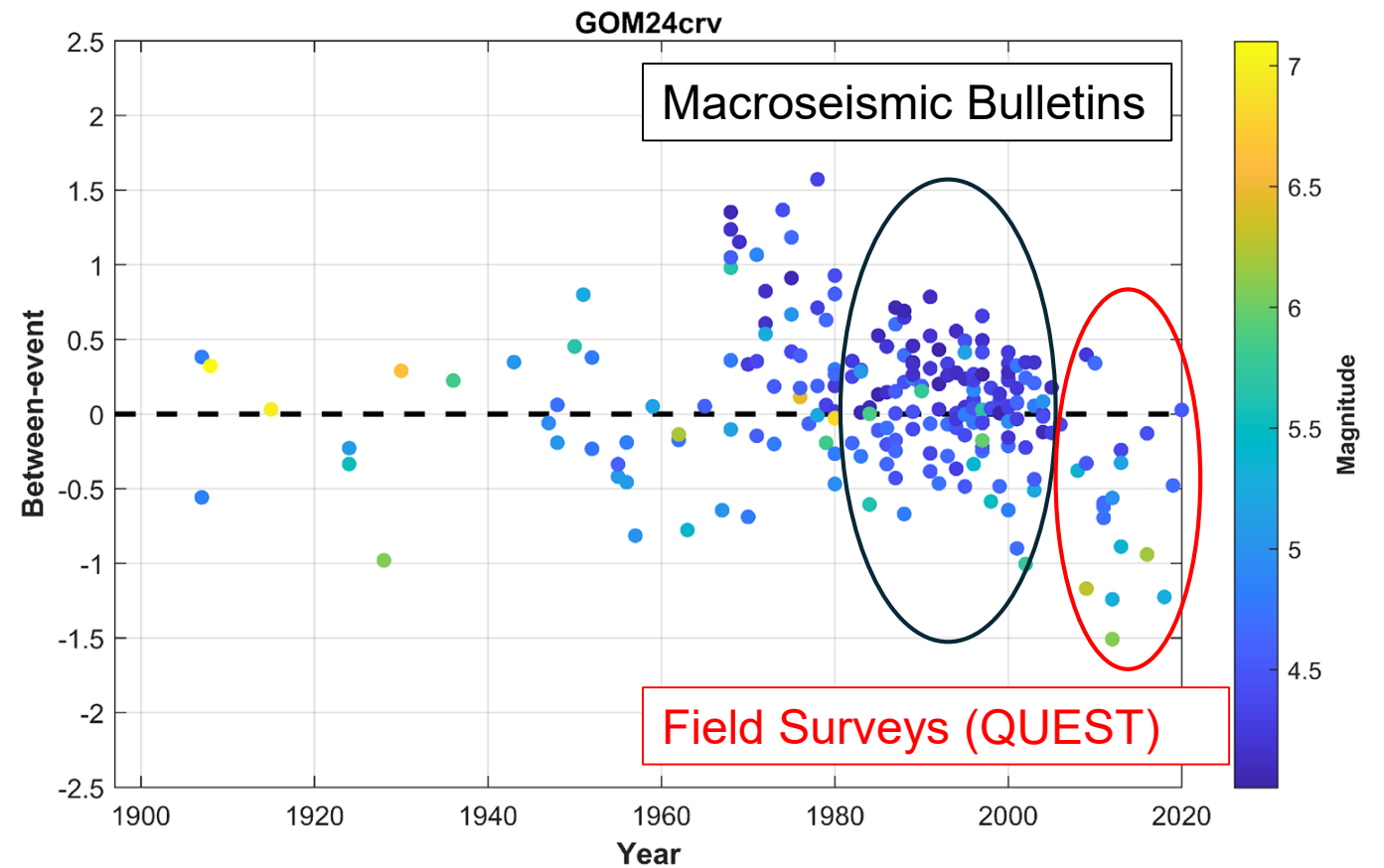
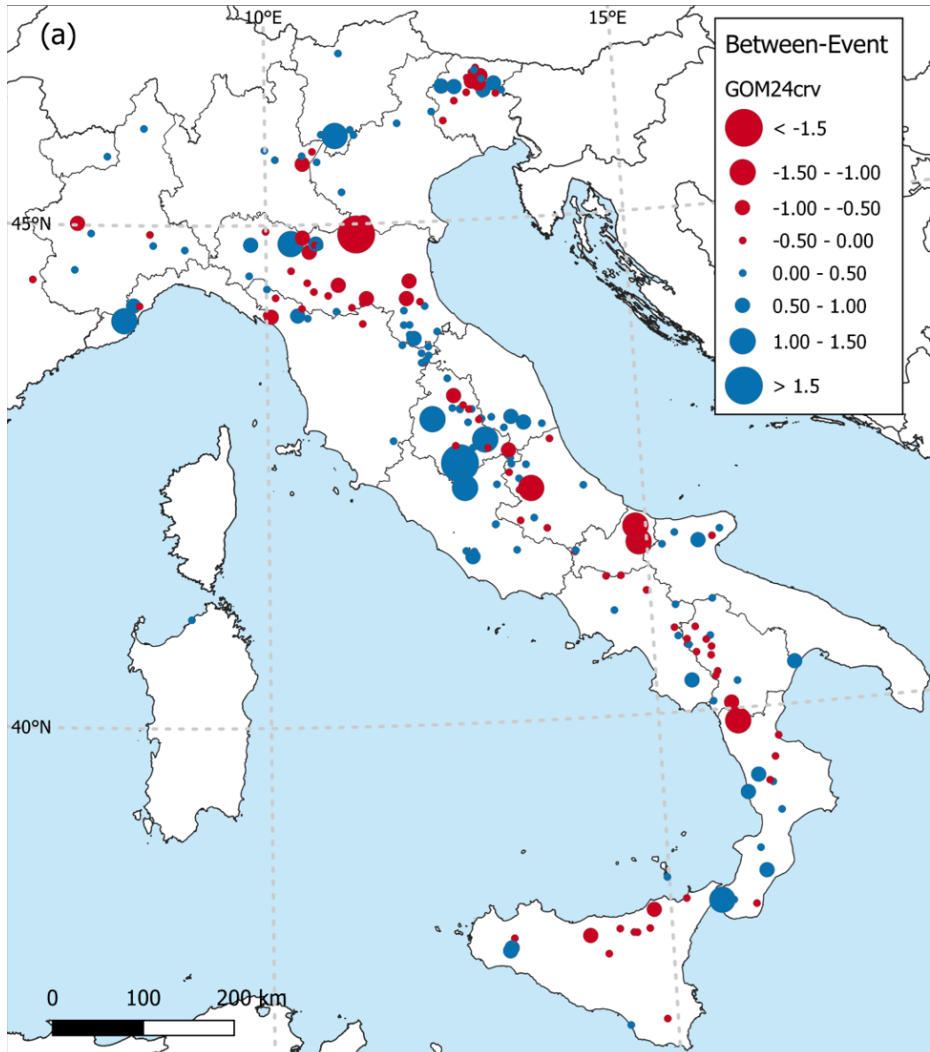
We assigned 5 points to the best-performing IPE (LLH or MAE values closest to zero) and 1 point to the worst

Model	LLH scoring				MAE scoring				Total score
	D #1	D #2	D #1alt	D #2alt	D #1	D #2	D #1alt	D #2alt	
ROV20	3	2	5	4	3	3	5	3	28
LOL24all	1	3	3	5	1	2	4	5	24
LOL24ins	2	1	2	3	2	1	2	2	15
GOM24	4	4	4	1	4	4	1	1	23
GOM24crv	5	5	1	2	5	5	3	4	30

4) Best-performing model

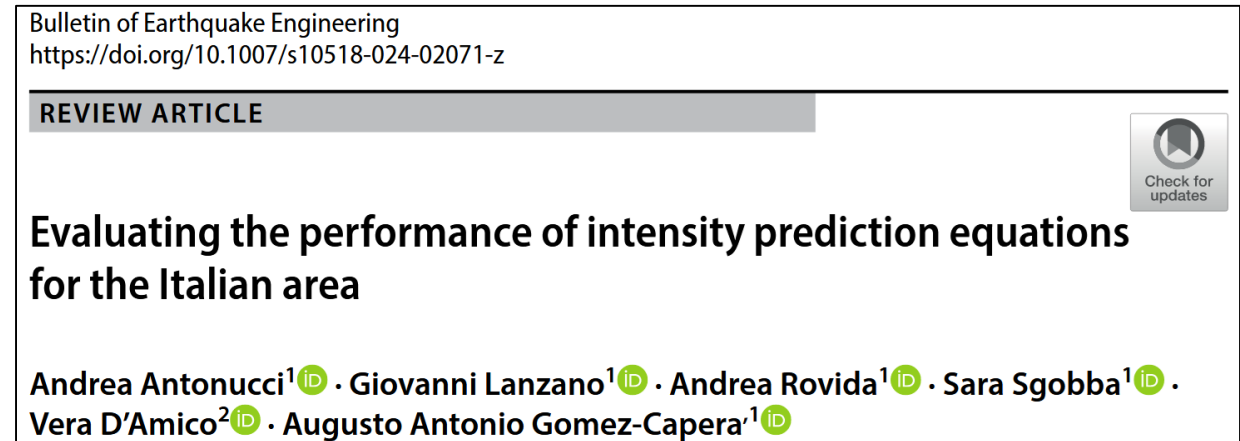
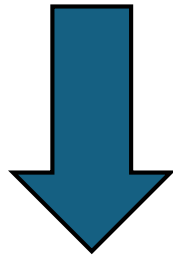
We estimated the **between-event** for each event as the mean of the residuals using GOM24crv

$$\delta B_e = \frac{1}{N_e} \sum_{i=1}^{N_e} R_{es}$$



Results and perspectives

- The GOM24crv model is **the best-performing model**
- The statistical indexes showed small differences among all models, indicating the validity of all the tested IPEs. These differences can be attributed to the **different criteria adopted for the selection of the calibration dataset of each model**



This study provided some hints:

- Incorporation of different distance metrics** (e.g., Joyner–Boore) in the prediction models
- Investigating the within-event terms (δW_{es})** to highlight the influence of rupture-directivity effects on a model prediction

Decomposition of residuals

We are reproducing the approach adopted in the GMMs to **decompose the residuals** into systematic effects to isolate of different sources of the model variability (Sgobba et al., 2021; Colavitti et al., 2022)

$$\mathbf{R}_{es} = \delta_0 + \delta B_e | \text{event} + \delta S2S_l | \text{locality} + \delta W_0$$

- δ_0 is the offset (approximately zero)
- δB_e is the between-event term (mean error of each event e)
- $\delta S2S_l$ is the site-to-site term (mean error for each locality l)
- **δW_0 is the aleatory error that contains unmodeled effects, like directivity**



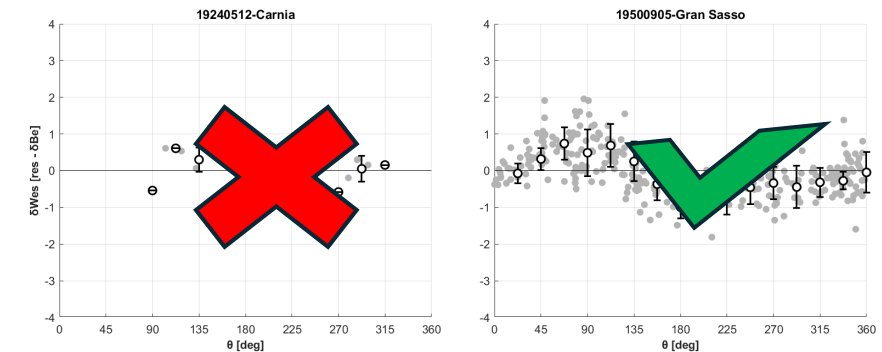
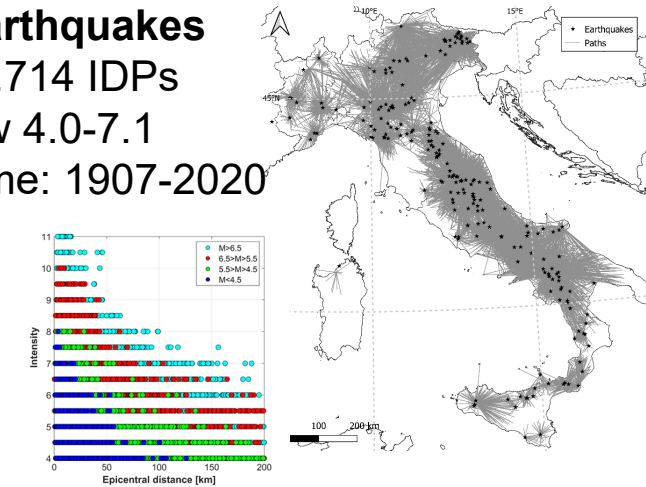
We are investigating the azimuthal dependence of the residuals (δW_0) to recognize potential directivity effects

Method

1. We used the event included in the **Dataset #1** (with instrumental location and magnitude)
2. We removed the events with **incomplete azimuthal sampling** through spatial analyses of macroseismic data
3. We **decomposed the residuals** at a single site for each event
4. We **analyzed the results in terms of δW_0** as a function of the fault geometry

213 earthquakes

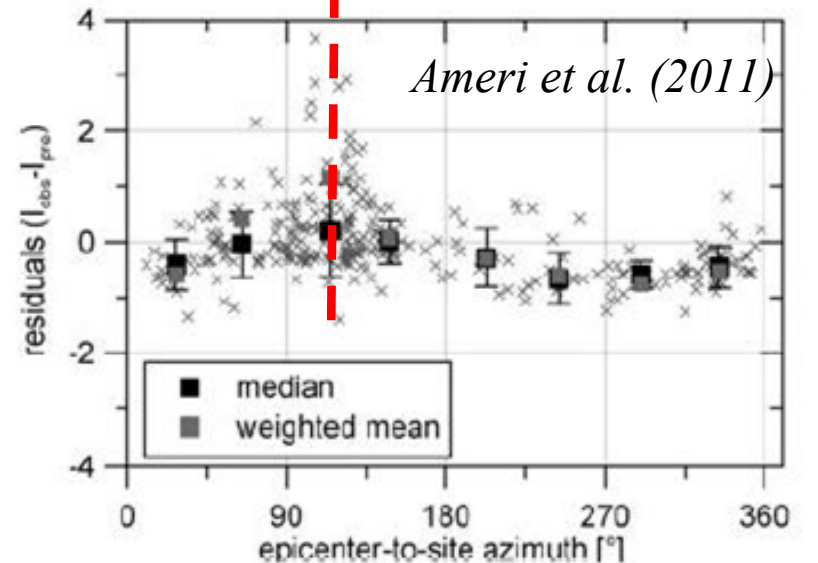
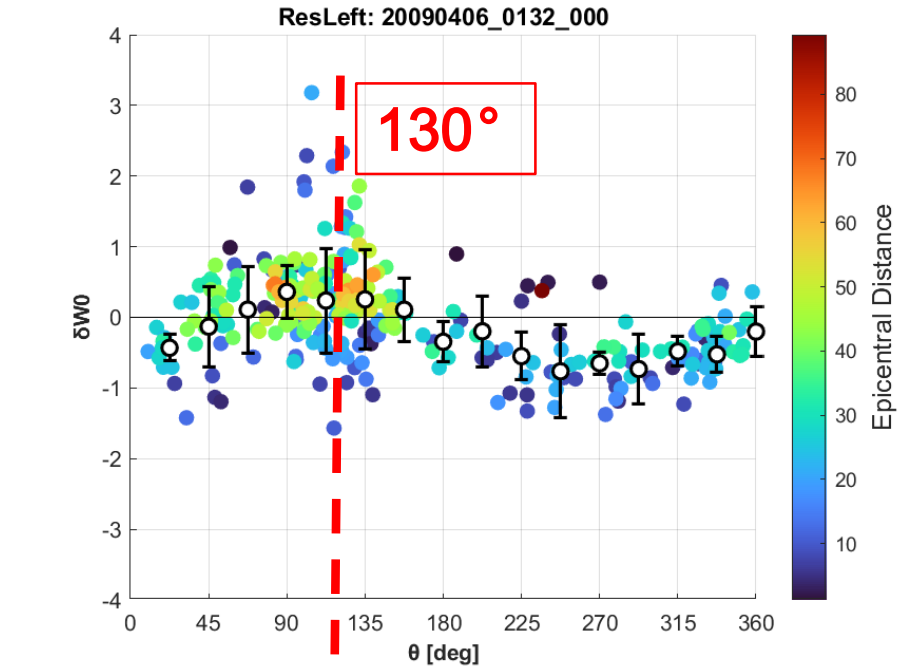
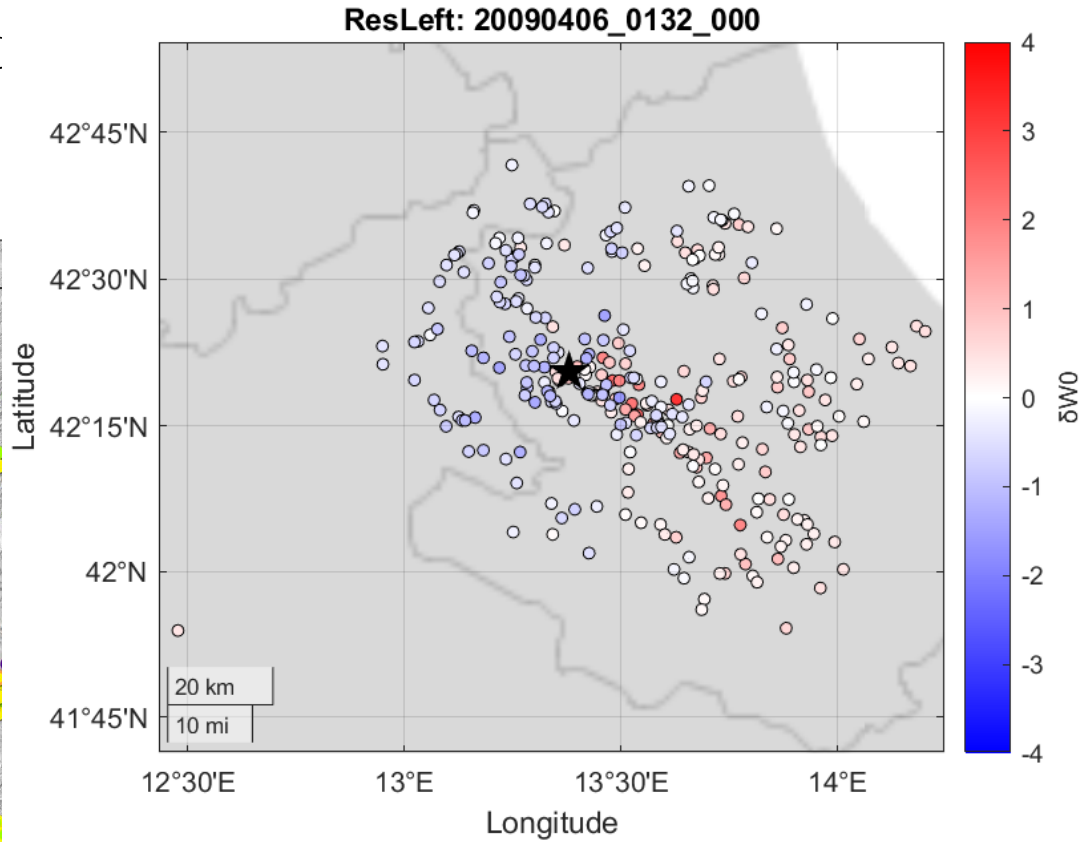
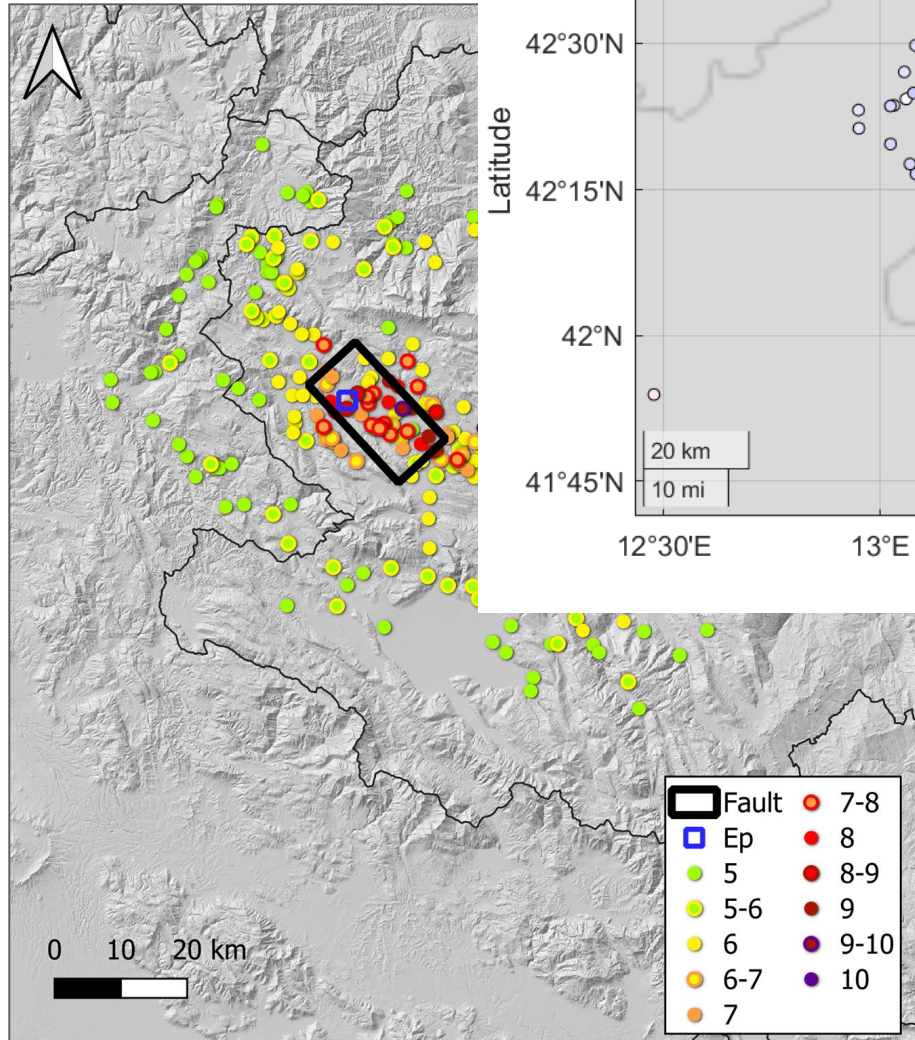
- 18.714 IDPs
- Mw 4.0-7.1
- Time: 1907-2020



$$R_{es} = \delta_0 + \delta B_e | \text{event} + \delta S2S_l | \text{locality} + \delta W_0$$

2009 - L'Aquila (Mw 6.3)

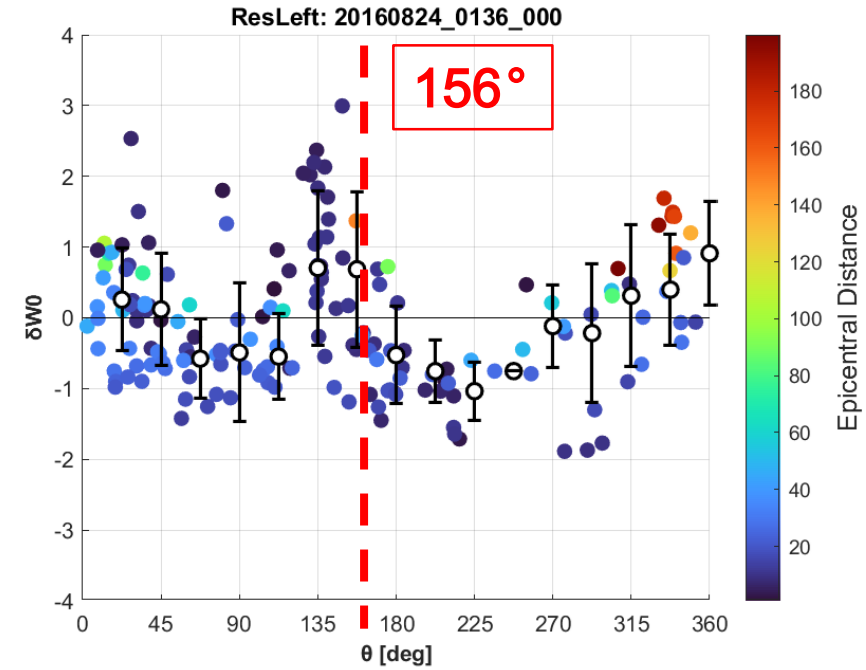
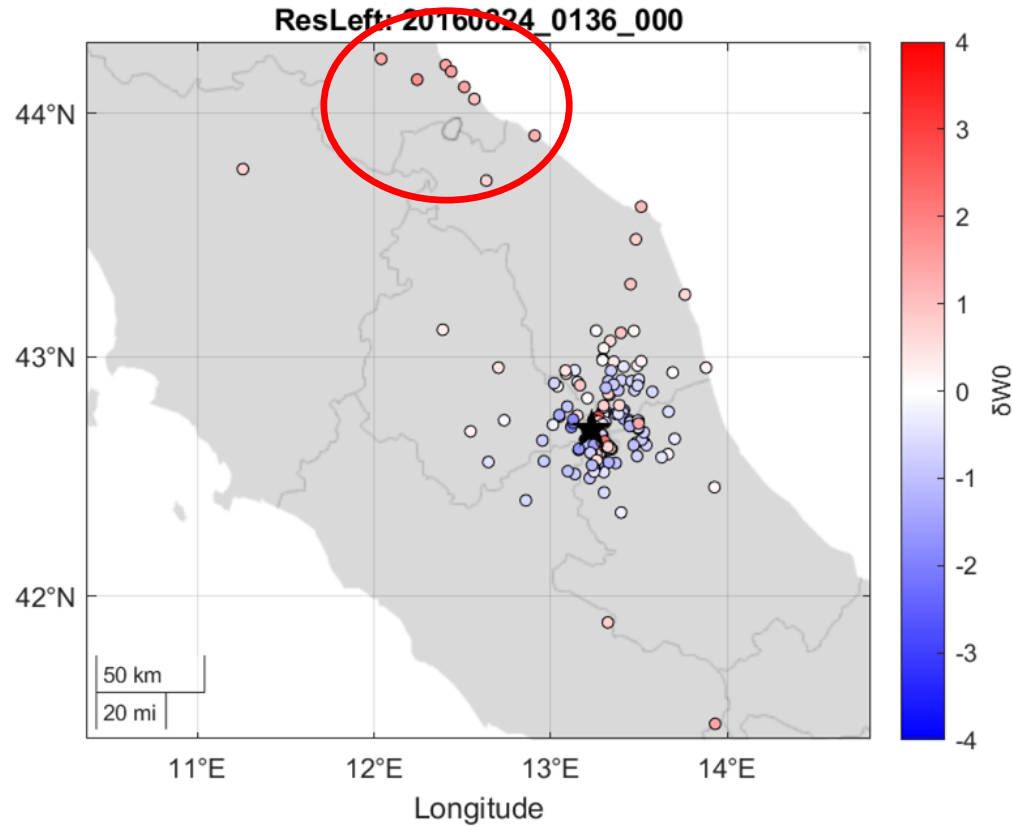
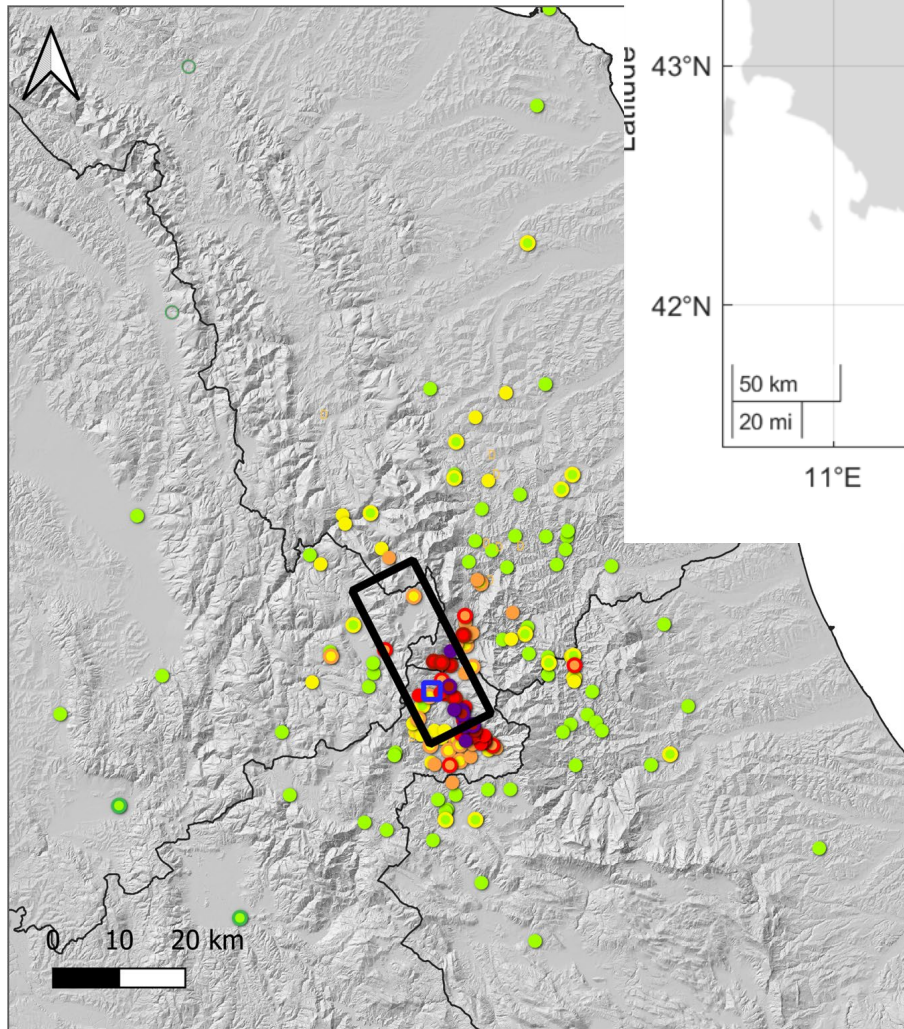
316 MDPs (Galli and C
Ix = 9-10



Ameri (2009)
Strike 140.0
Dip 50.0
Rake -90.0

2016 - Amatrice (Mw 6.1)

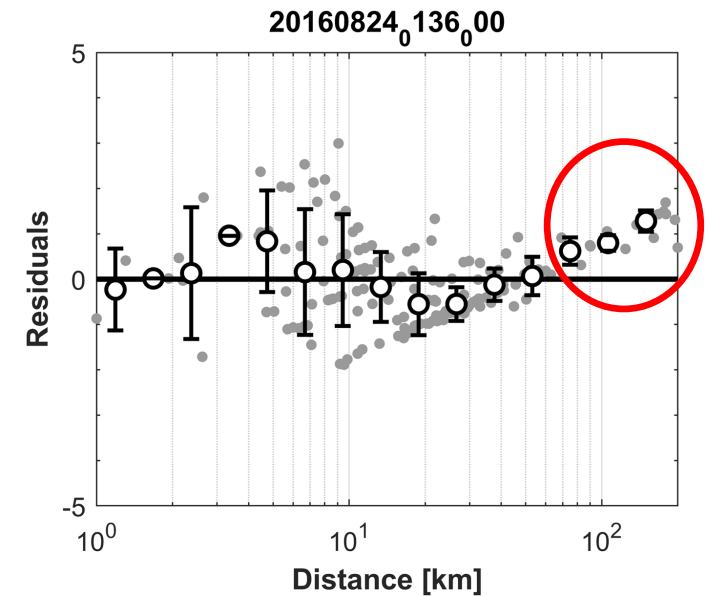
221 MDPs (Rossi et al, 2019)
 $I_x = 10$

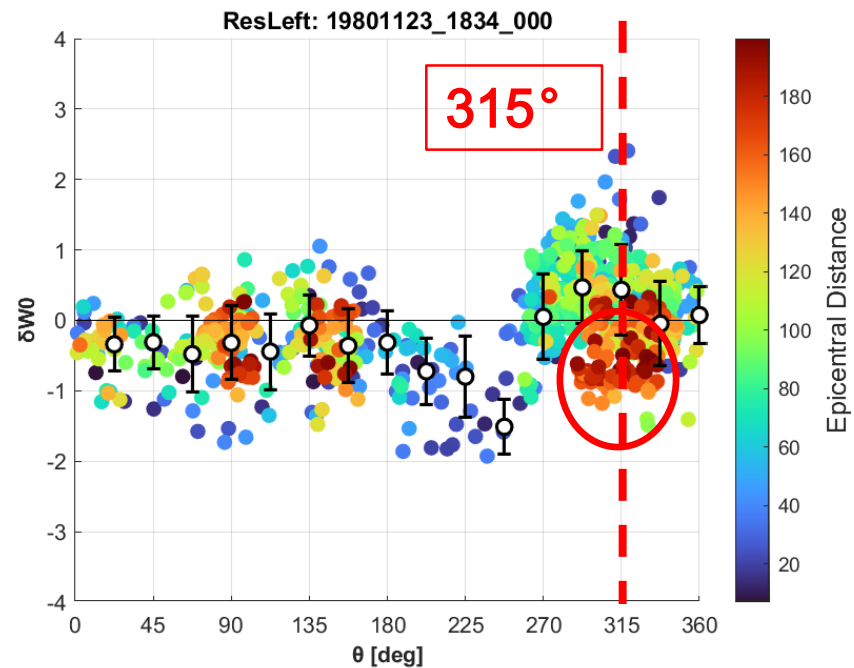


Complex azimuthal
distribution

Tinti (2016)
Strike 156.0
Dip 50.0
Rake -85.0

**Positive Bias for
long distance
data ($R > 80\text{km}$)**





Mw 6.8 (CPTI15)

Ix 10

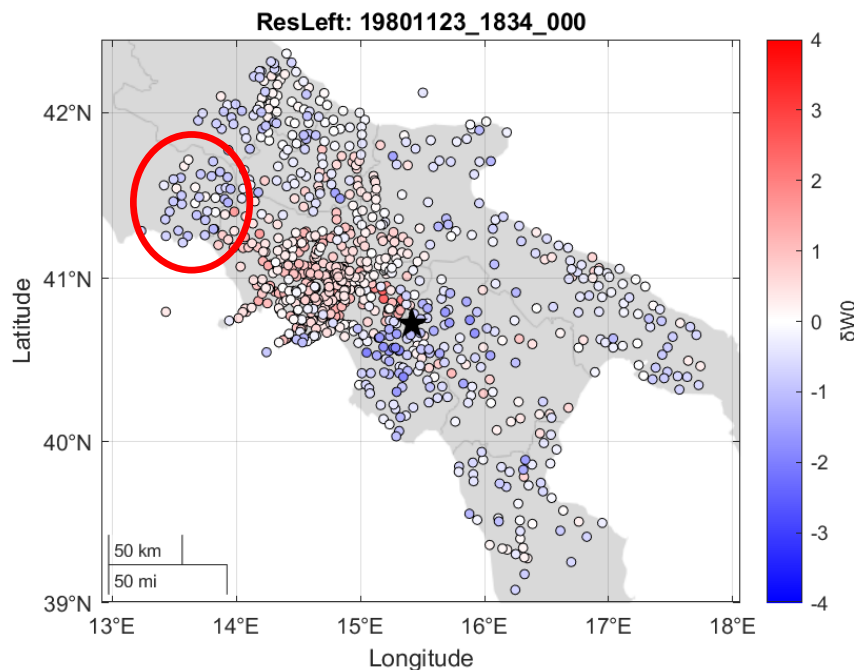
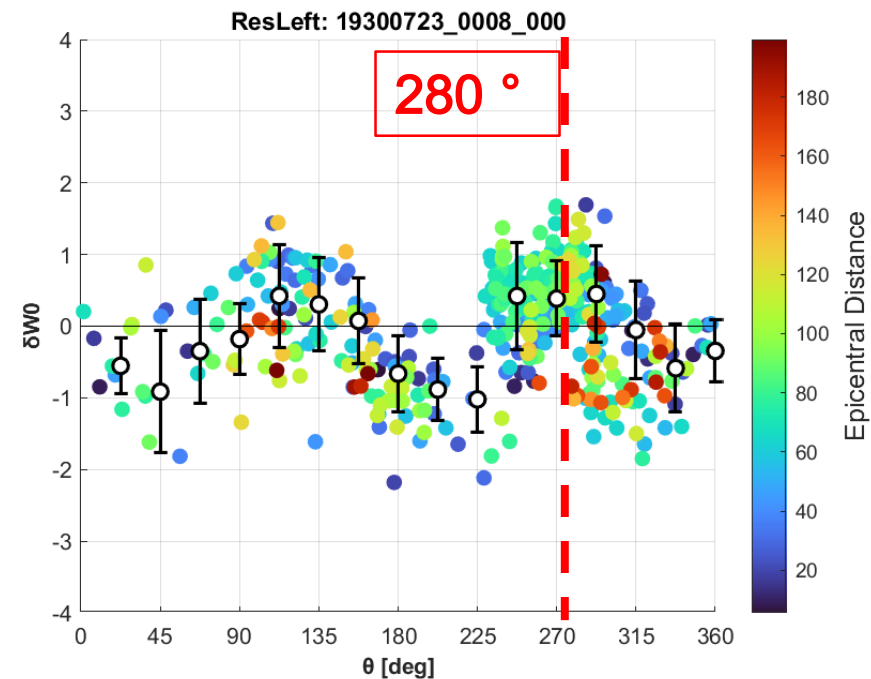
Ameri et al. (2011)

Strike 315

Dip 60

Rake -90

Negative
residuals at
long distances
(>120 km)



Mw 6.6 (CPTI15)

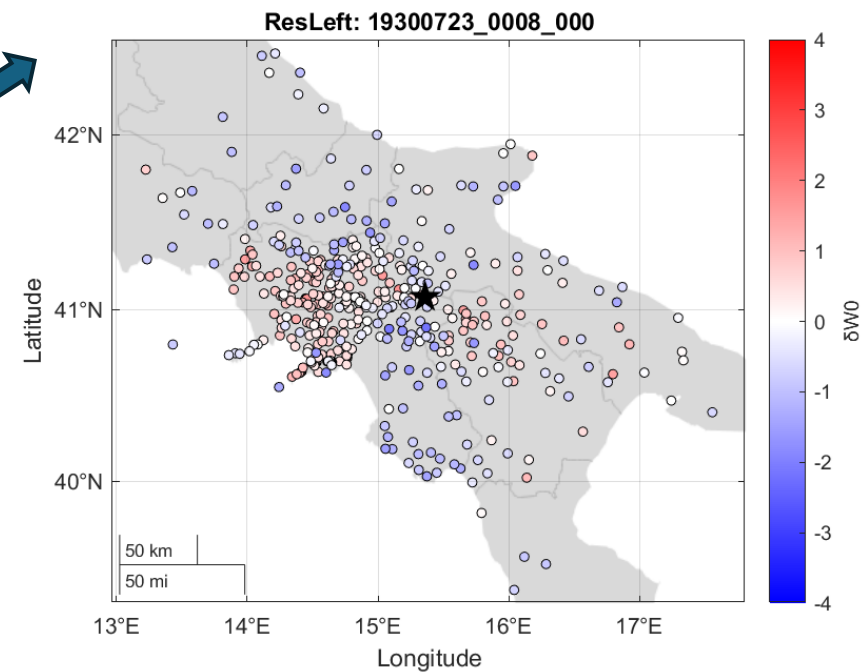
Ix 10

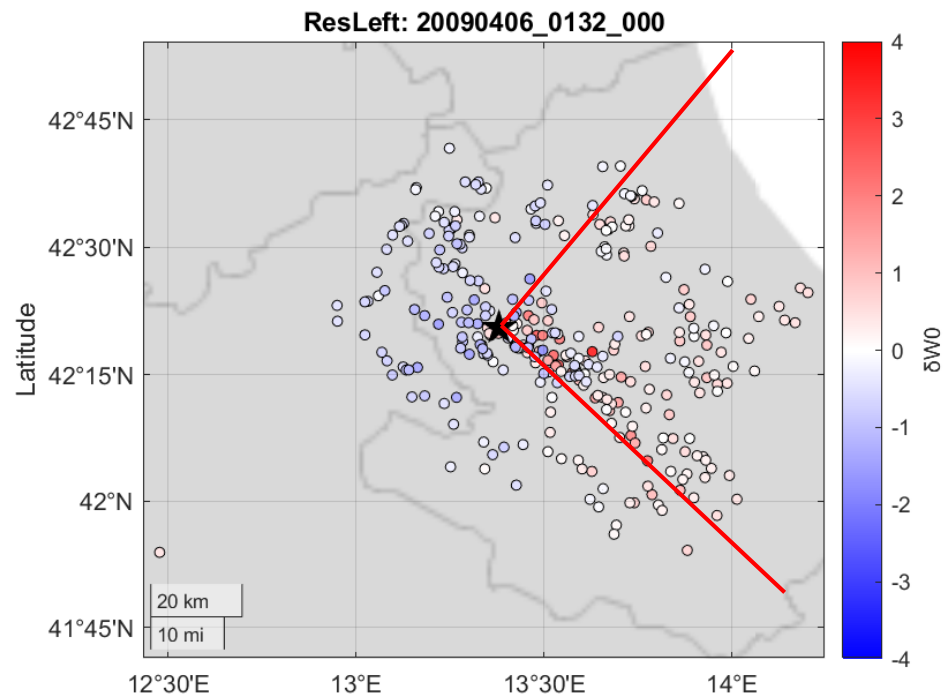
DISS - ITIS088

Strike 280

Dip 64

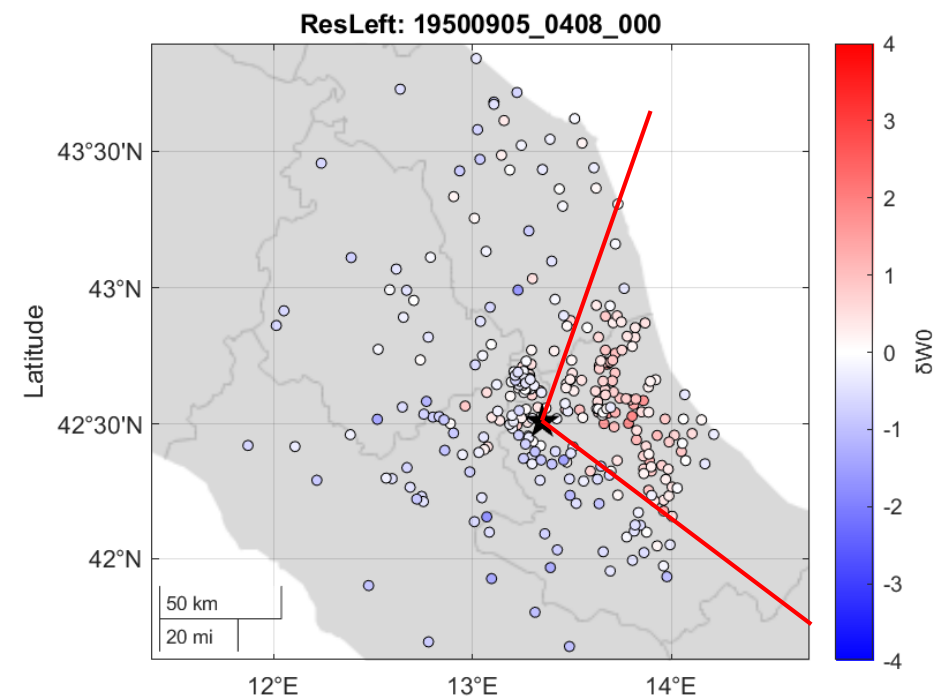
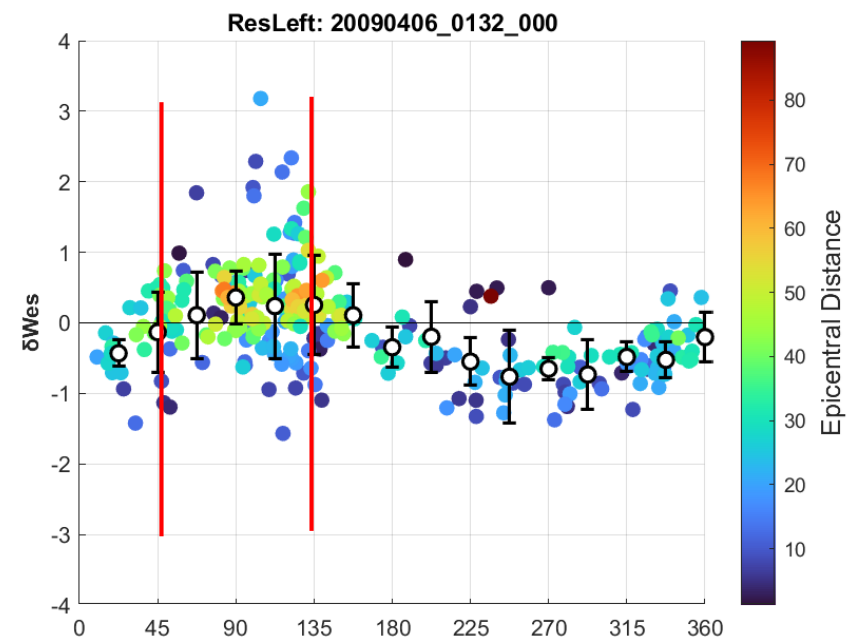
Rake 237





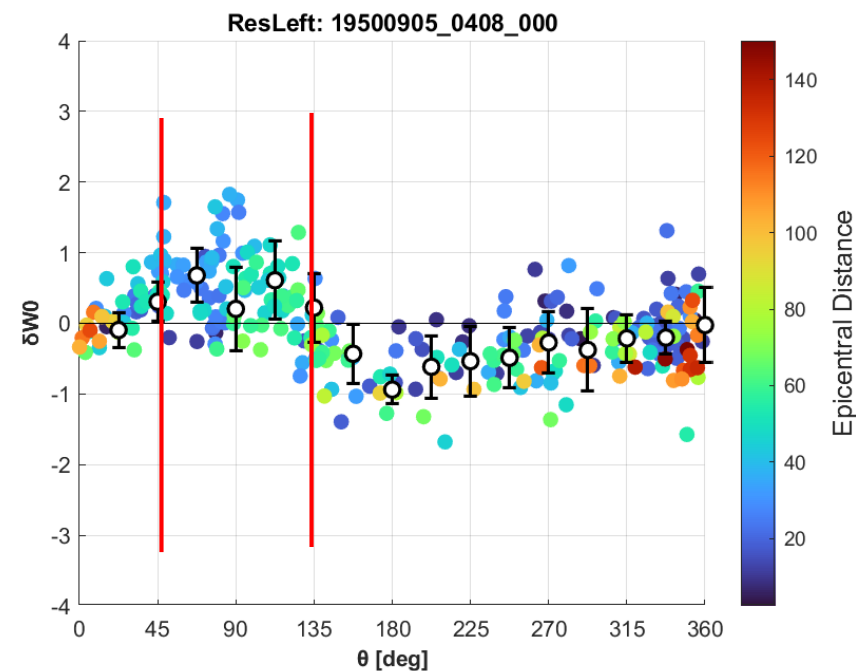
2009-04-06
Mw 6.3 (CPTI15)

18 km of distance
from 2009 – L'Aquila



1950-09-05
Mw 5.7 (CPTI15)

Directivity effects at
sites located
between 45° and 135°



Preliminary results and next steps

- We are carrying out an exploratory **systematic analysis to detect directivity effects in the macroseismic data through the approach based on the residuals**, originally developed for ground motion parameters obtained from the recordings
- We preliminary observed **positive residuals** ($I_{\text{obs}} > I_{\text{pred}}$) in a given direction **corresponding to strike of the fault**



Extending the dataset to past events with macroseismic magnitudes and locations

Thank you

Andrea Antonucci

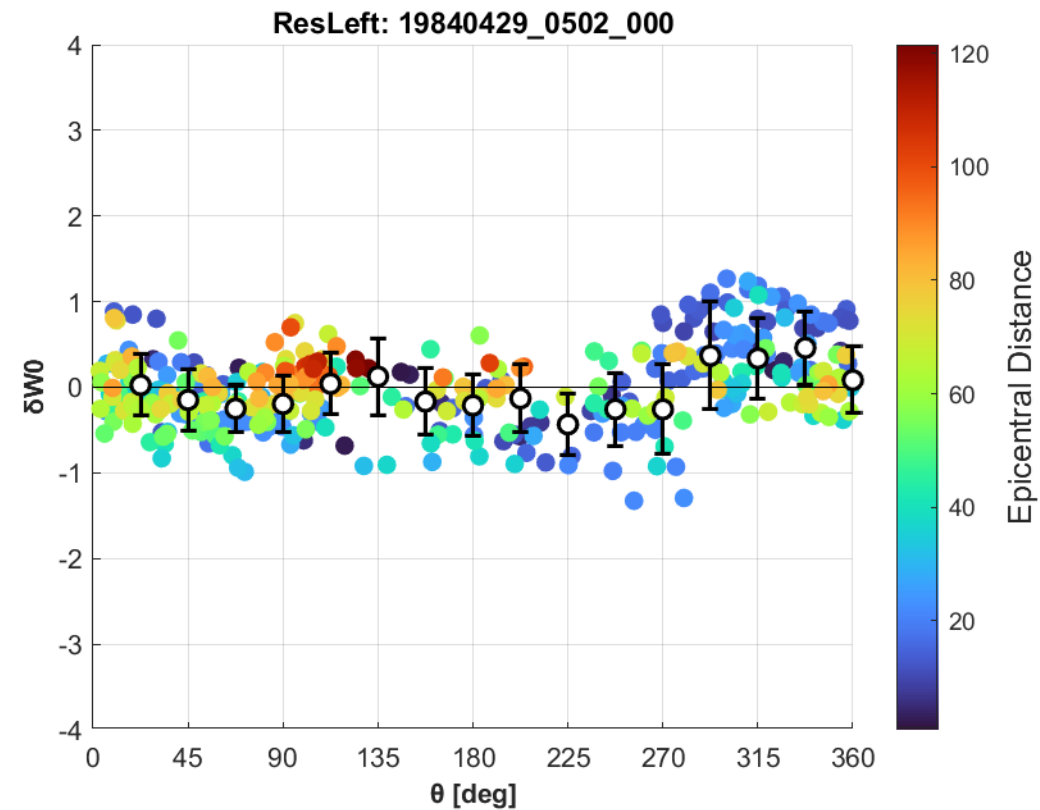
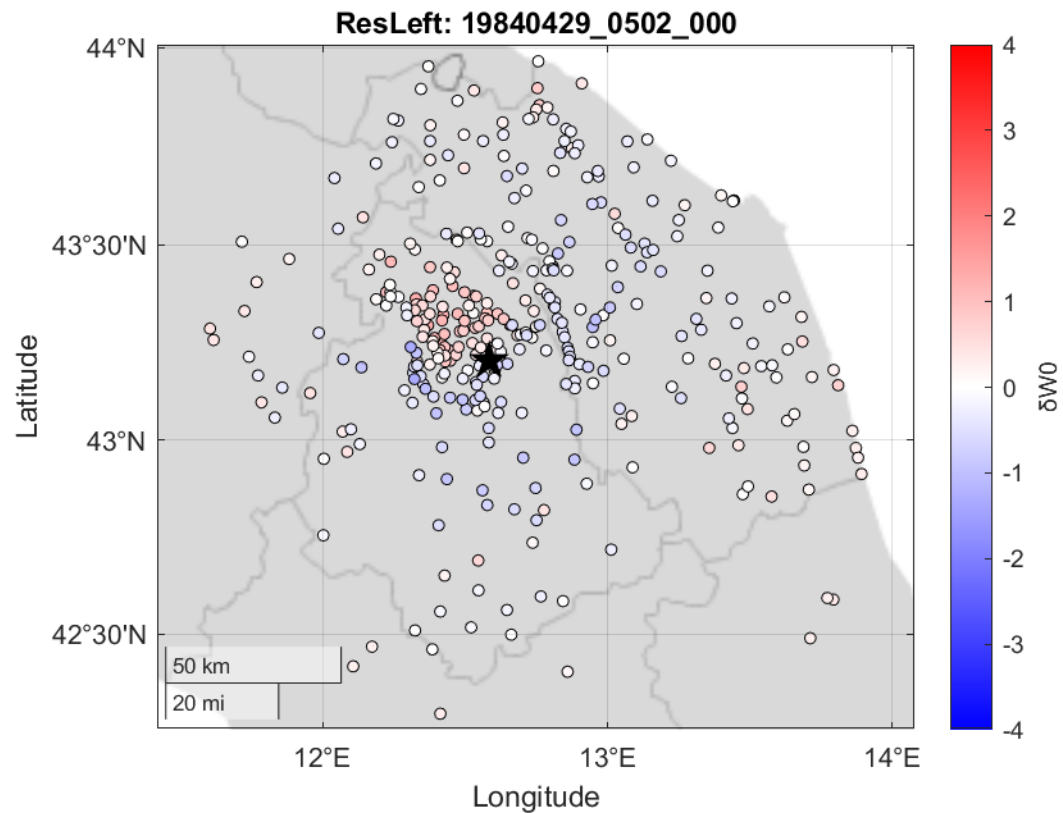
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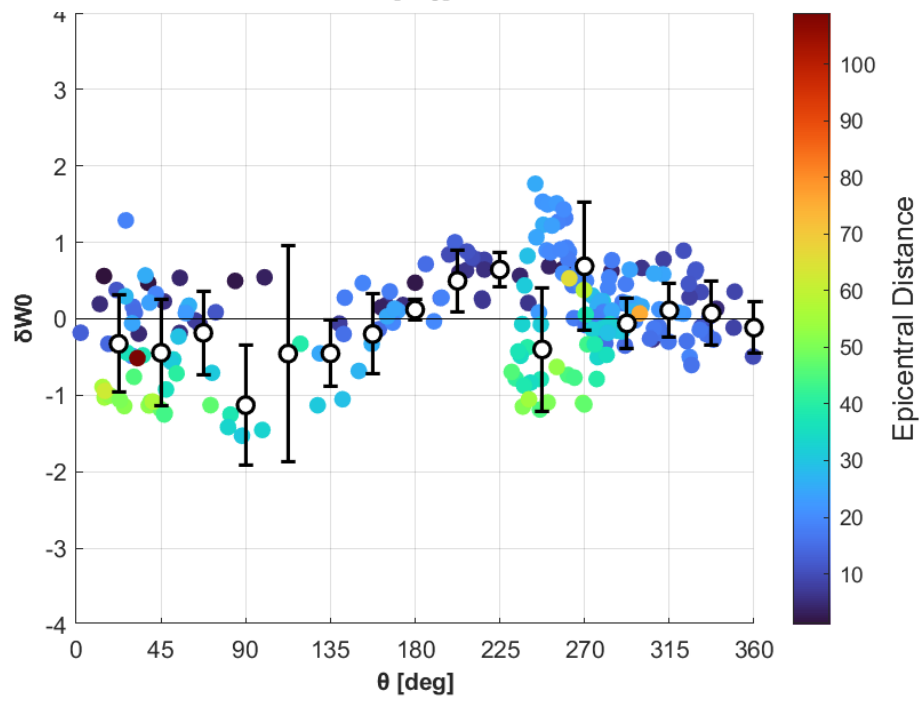
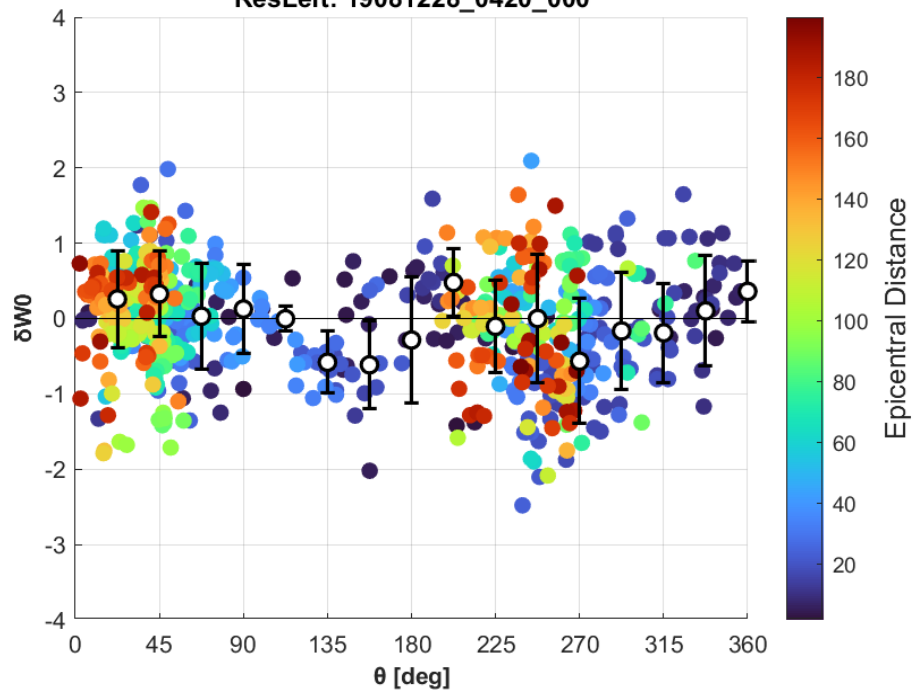


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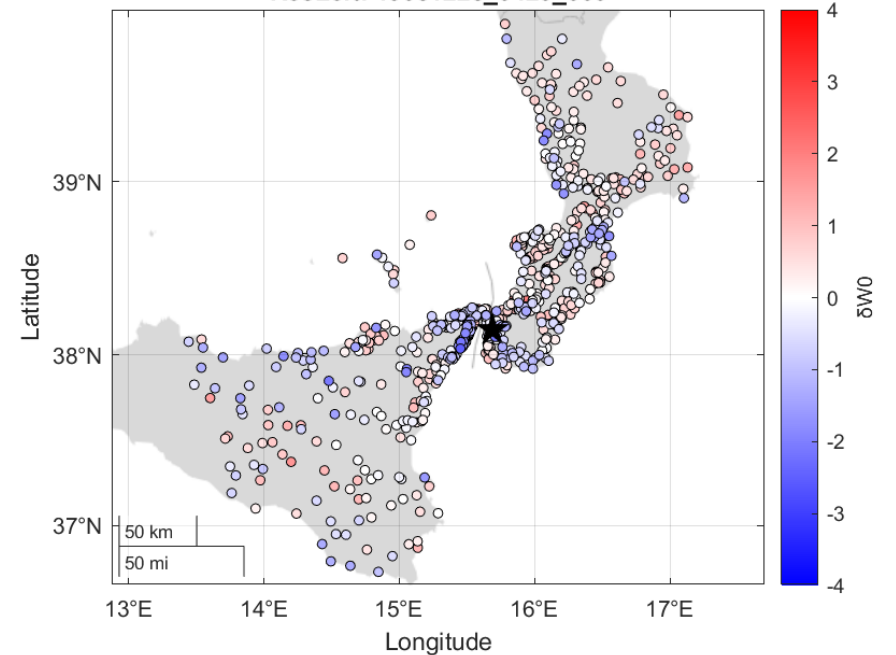
Ameri et al (2009)
Strike 130°
Dip 20°
Rake -90°

ResLeft: 19081228_0420_000

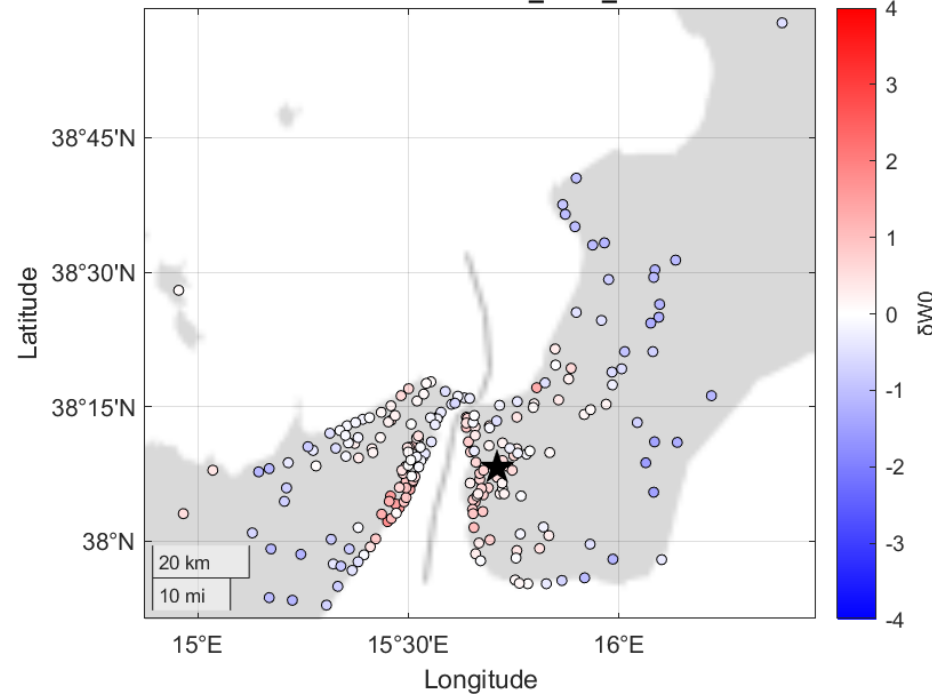


Faglia del 1908
Secondo Natale
and Pingue
(1991) ha
provocato evento
1975

ResLeft: 19081228_0420_000

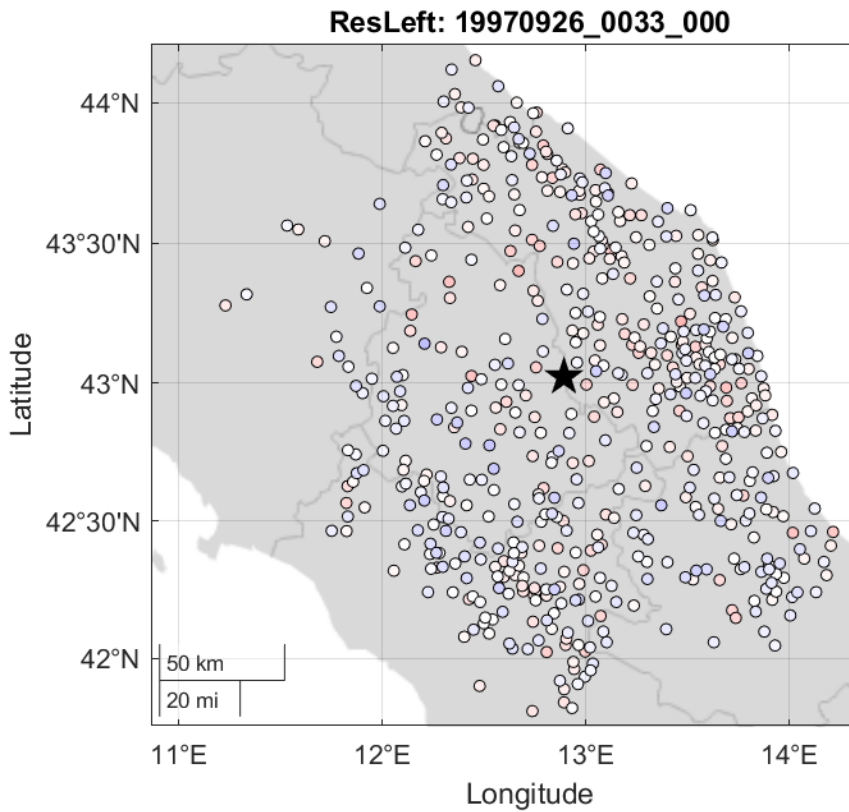


ResLeft: 19750116_0009_000

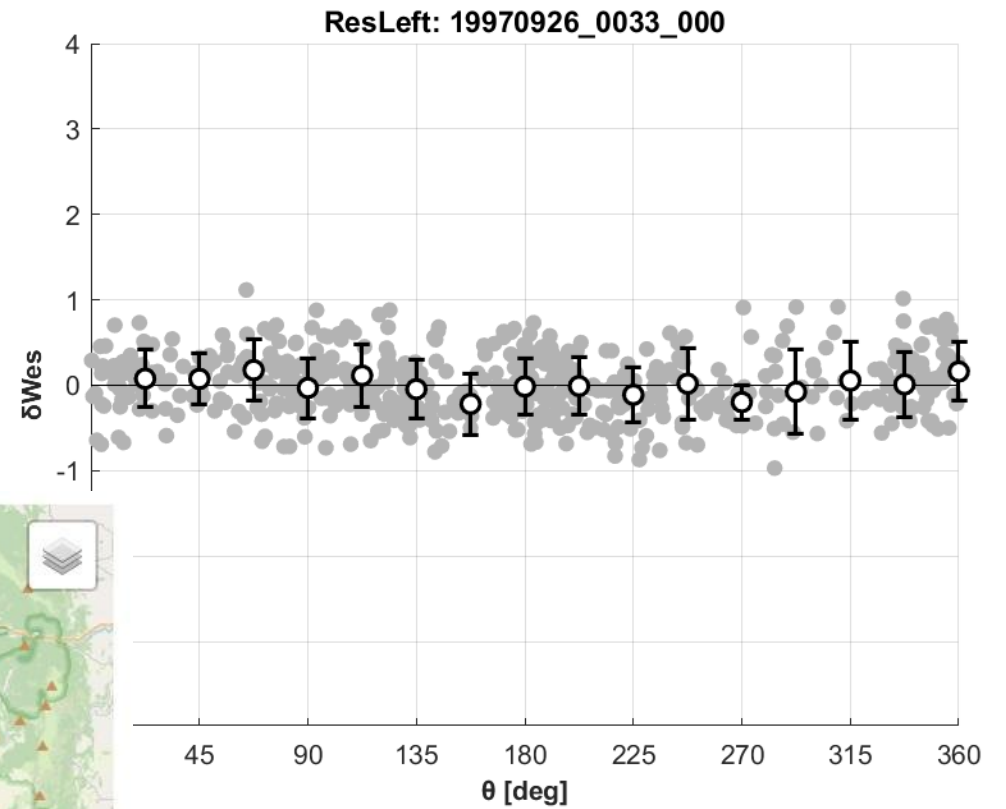


Results examples: Umbria-Marche (1997)

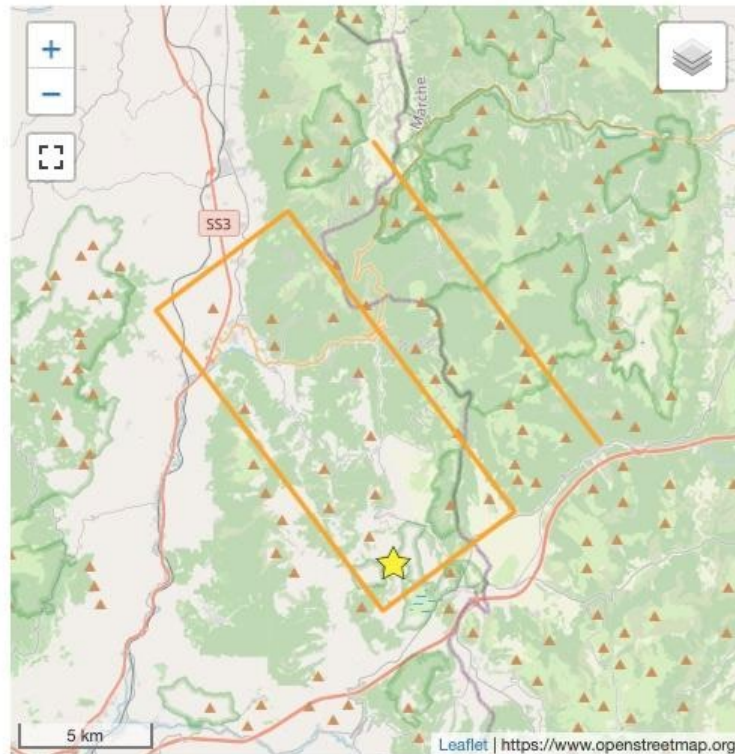
March 6



No signature of
source
directivity



No residual bias
with distance



DISS
Strike143.0
Dip40.0
Rake-85.0